

Some Climatic Variability on the Argentina Regions
Walter M.Vargas

CONICET – UBA
ARGENTINA

The following presentation is a sample of the lines of work performed by an investigation group which respond to the following objectives:

Climatology of precipitation, temperature and streamflow extremes.

Low variability (climatic jumps, trends) .

Synoptic Climatology of precipitation and temperature. Relationship between climatic and hydrologic variables.

In an attempt to explain climatic jumps and variabilities observed in precipitation series and water vapor advection over the southern cone of South America, the estimated low frequencies of pressure series at the center of south Atlantic anticyclone are presented and compared with data from continental coastal stations (Buenos Aires Observatory and Rio de Janeiro). In figure 1 we can distinguish a low frequency wave in phase and therefore certain properties of the center of the anticyclone can be inferred from the pressure series over the continent. This is also related to the Río Paraná streamflow and rainfall in northwestern Argentina. These variabilities are accompanied by climatic jumps in annual precipitation which were found for the years shown in figure 2 over the Argentine territory. Nevertheless this effect is non-existent in Chile, where the annual precipitation series only present trends . It is important to point out that these jumps coincide with important changes in the economic activities of the region (specially agriculture). These changes can be inferred from the progressive expansion of cultivated areas towards semi arid or arid regions, due to the lapse of above normal precipitations in the region caused by the climatic jump. One of the regions of the Rio de la Plata Basin with greatest flood risk and considerable impacts is the Uruguay river basin. In this area the necessity of a hydrologic forecast requires the study of precipitation properties and their relationship with streamflows. First of all, the basin is affected by different precipitation regimes which are clearly defined by the annual cycle. It can be inferred, through the mean area precipitation and streamflow in the gauge stations which limit the runoff area in each subregion, different annual patterns as those in figure 3. In the case of the northern region, called R1, streamflows are total values and not differences between stations as in the other Rs. It is clear that a precipitation – streamflow model for the Uruguay basin there, should contemplate the different relationships between the cited variables determined by the soil in each sub – basin. Another element of interest and which reflects the kind of relation between the mentioned variables and its stability is given by the laws of probability associated with the anomalies. In this case, it is of interest to compare the amount of negative anomalies each year and the sequences of those properties. In this sense a binomial and a geometric law are proposed and the precipitation and streamflow series at selected points of the basin are fitted. The stations chosen for fitting monthly precipitation are Rio de Janeiro and Concordia, and Santo Tome and Concordia are selected in the streamflow adjustment. In both cases it is intended to show the conditions in the northern and southern limits of the basin. On the other hand, this basin presents the peculiarity of

being very extended in latitude when compared to its maximum zonal extension. Tables 1 show the respective fittings and it can be inferred that precipitation fits very well throughout the basin indicating similar processes in the climatic scale, which are not observed in the streamflow, which although statistically homogeneous along the river, do not present the same fitting models in the case of negative monthly anomalies. This is due, mainly, to two facts: persistence is greater in streamflow and double maximums are generated in the number of monthly anomalies each year. This indicate that the physical relation between precipitation and streamflow is not linear.

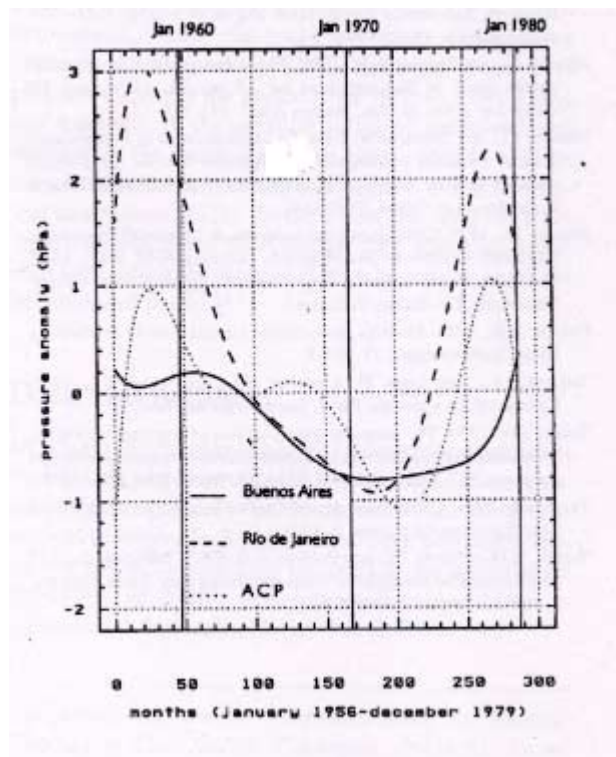


Fig. 1: Pressure anomaly series of the Anticyclone center (ACP), Buenos Aires and Rio de Janeiro.

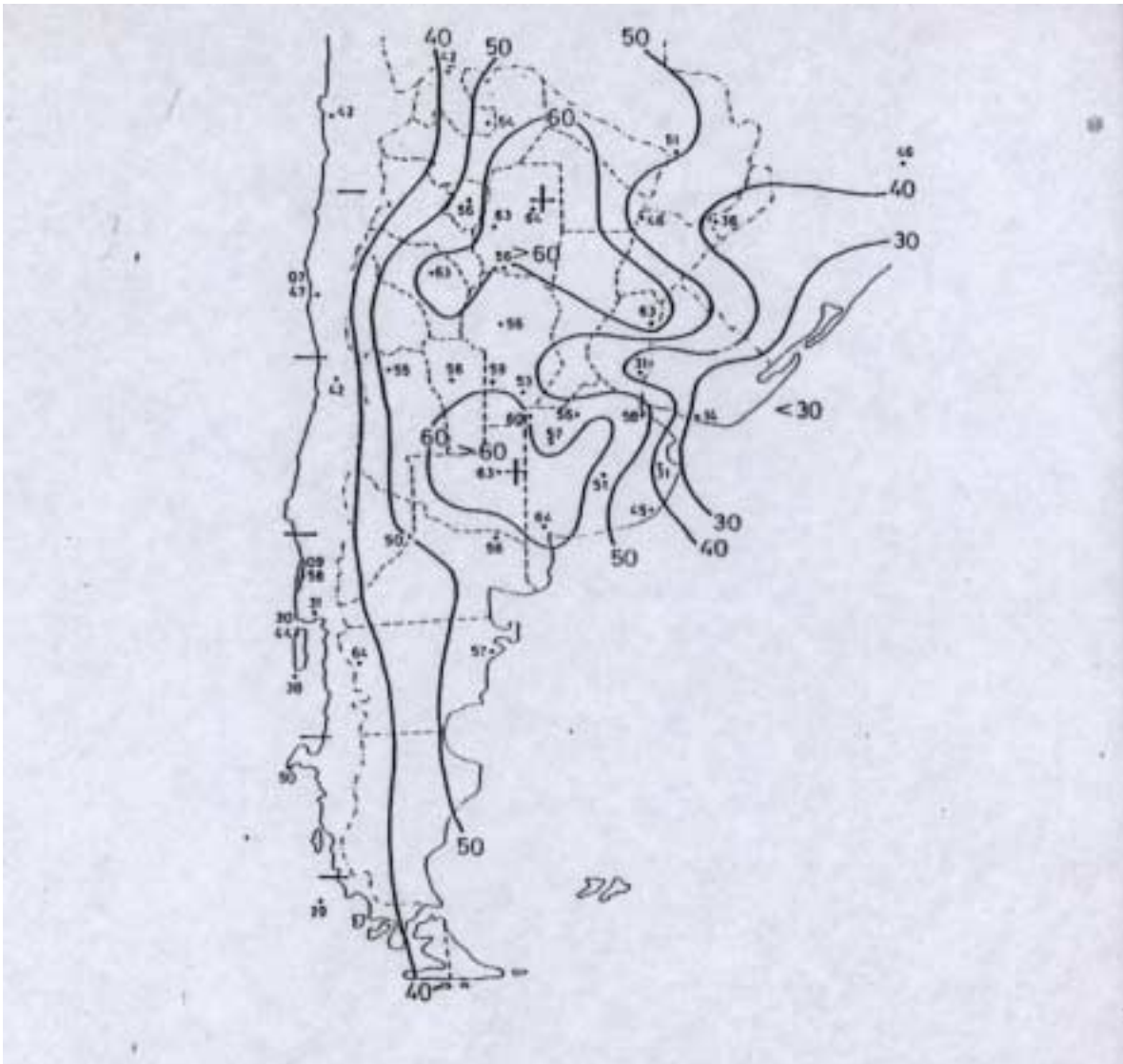


Fig. 2 Isochrons of years registering the main change in the mean of annual precipitation. Each number must be read plus 1900.

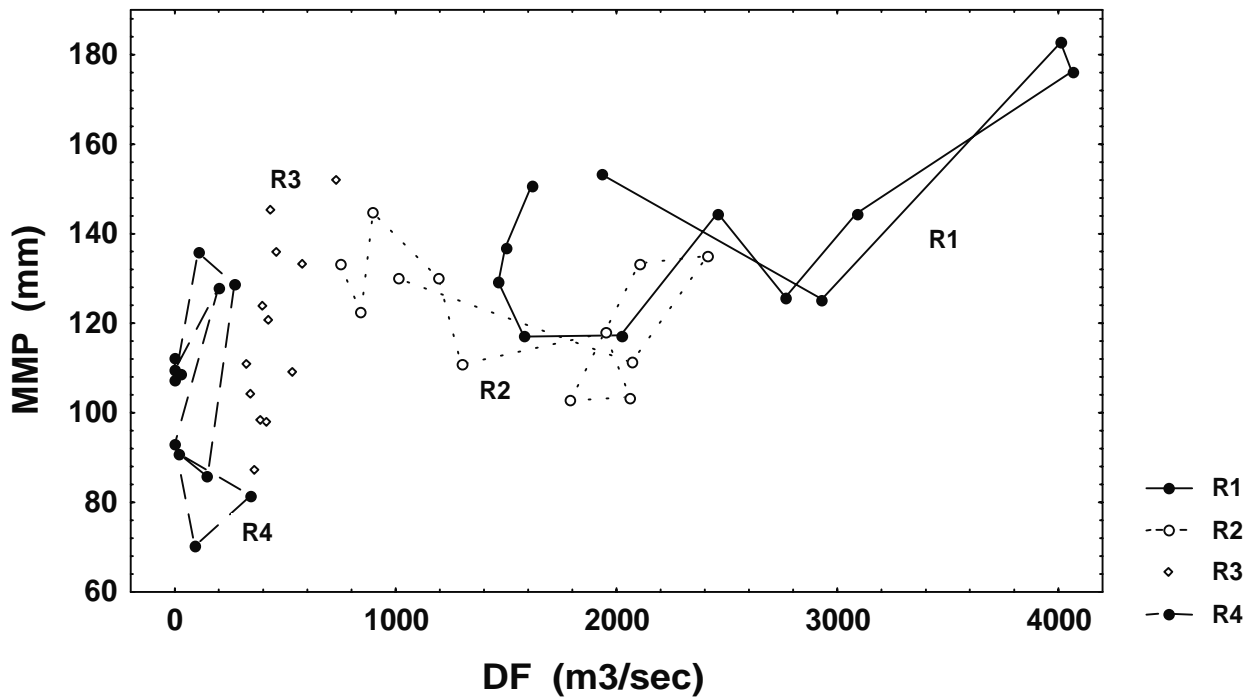


Fig 3. Annual monthly means precipitation (MMP) vs discharges (DF) at four basins of the Uruguay River

NEGATIVE ANOMALIES
PRECIPITATION

a) BINOMIAL DISTRIBUTION
CONCORDIA

RIO DE JANEIRO

K	F _k	PROB. ESTIMATE	FREQ. THEORETICAL	K	FK	PROB. ESTIMATE	FREQ. THEORETICAL
0	1	.000027	0.0	0	0	.000037	0.0
1	0	.000449	0.0	1	0	.000592	0.0
2	1	.003474	0.3	2	1	.004367	0.5
3	0	.016277	1.4	3	5	.019534	2.6
4	6	.051470	4.5	4	8	.058977	8.0
5	9	.115737	10.3	5	20	.126625	17.3
6	19	.189768	16.8	6	23	.198234	27.1
7	17	.228600	20.3	7	27	.228005	31.2
8	15	.200798	17.8	8	27	.191223	26.1
9	13	.125423	11.1	9	17	.114043	15.6
10	5	.052881	4.7	10	4	.045910	6.2
11	3	.013513	1.2	11	3	.011201	1.5
12	0	.001583	0.1	12	2	.001253	0.1

$\chi^2 = 2.781$

$\chi^2 = 3.96$

**b) GEOMETRIC DISTRIBUTION
CONCORDIA**

RIO DE JANEIRO

K	F	PROBABILITY ESTIMATE	FREQUENCY THEORETICAL	K	F	PROBABILITY ESTIMATE	FREQUENCY THEORETICAL
0	387	.592453	1372.0	0	561	.56962	538.2
1	140	.241452	151.6	1	215	.24515	231.6
2	54	.098404	61.7	2	92	.10551	99.7
3	24	.040104	25.1	3	41	.04541	42.9
4	14	.016344	10.2	4	14	.01954	18.4
5	3	.006661	4.1	5	11	.00841	7.9
6	2	.002715	1.7	6	2	.00362	3.4
7	3	.001106	0.6	7	4	.00156	1.4
8	1	.000451	0.2	8	4	.00070	0.6
9	0	.000184	0.1	9	1	.00028	0.2
10	0	.000075	0.0	10	0	.00012	0.1
11	0	.000031	0.0	11	0	.00005	0.0
12	0	.000000	0.0	12	0	.00000	0.0

$\chi^2 = 7.65$

$\chi^2 = 22.03$

**NEGATIVE ANOMALIES
RUNOFF**

**a) BINOMIAL DISTRIBUTION
CONCORDIA**

SANTO TOME

K	Fk	PROB. ESTIMATE	FREQ. THEORETICAL	K	FK	PROB. ESTIMATE	FREQ. THEORETICAL
0	0	.000015	0.0	0	0	.000010	0.0
1	1	.000273	0.0	1	0	.000193	0.1
2	4	.002286	0.2	2	2	.001712	0.1
3	3	.011619	1.1	3	2	.009186	0.7
4	13	.039863	3.8	4	8	.033277	2.7
5	10	.097269	9.4	5	8	.085727	7.0
6	6	.173051	16.7	6	10	.161032	13.2
7	7	.226194	21.9	7	16	.222235	18.2
8	18	.215584	20.9	8	7	.223635	18.3
9	16	.146113	14.1	9	10	.160031	13.1
10	5	.066844	6.48	10	7	.077299	6.3
11	8	.018533	1.7	11	7	.022629	1.8
12	6	.002355	0.2	12	5	.003036	0.2

$\chi^2 = 136.22$

$\chi^2 = 74.74$

**b) GEOMETRIC DISTRIBUTION
CONCORDIA**

SANTO TOME

K	F	PROBABILITY ESTIMATE	FREQUENCY THEORETICAL	K	F	PROBABILITY ESTIMATE	FREQUENCY THEORETICAL
0	289	.46326	210.3	0	249	.44877	180.8
1	57	.24865	112.8	1	57	.24738	99.6
2	27	.13346	60.5	2	24	.13636	54.9
3	27	.07163	32.5	3	13	.07516	30.2
4	13	.03845	17.4	4	19	.04143	16.7
5	11	.02064	9.3	5	16	.02284	9.2
6	8	.01108	5.0	6	9	.01259	5.0
7	8	.00594	2.7	7	6	.00694	2.7
8	8	.00319	1.4	8	0	.00382	1.5
9	2	.00171	0.7	9	3	.00211	0.8
10	3	.00092	0.4	10	5	.00116	0.4
11	1	.00049	0.2	11	2	.00064	0.2
12	0	.00	0.0	12	0	.00000	0.0

$\chi^2 = 132.46$

$\chi^2 = 98.49$

TABLES 1